

CONTRIBUTIONS TO THE THEORY  
OF ATMOSPHERIC REFRACTION (1)



Abstract

Since the barometer measures the weight of the overlying atmosphere, it follows by the law of Gladstone and Dale that the height integral  $\int (n - 1) dr$  of the atmospheric refractivity for light, taken from ground level up to the top of the atmosphere, is directly proportional to ground pressure. The refractivity integral, therefore, can be determined without detailed knowledge of the height distribution of the refractive index, which not only simplifies the derivation of refraction formulas in which atmospheric models have been used hitherto, but also improves their accuracy. For zenith distances not exceeding about 75 degrees, the correction for astronomical refraction will be given by the standard formula

$$\Delta z_0'' = 16''.271 \tan z \left[ 1 + 0.0000394 \tan^2 z \left( \frac{p - 0.156e}{T} \right) \right] \left( \frac{p - 0.156e}{T} \right) - 0''.0749 (\tan^3 z + \tan z) \left( \frac{p}{1000} \right)$$

where  $z$  is the apparent zenith distance,  $p$  is the total pressure and  $e$  is the partial pressure of water vapour, both in millibars, and  $T$  is the absolute temperature in degrees Kelvin. Part II of the paper contains further applications of the theory to refraction problems in satellite geodesy, including the photogrammetric refraction and the atmospheric corrections in the ranging of artificial satellites.

Part I. Astronomical Refraction

Derivation of General Formula for Astronomical Refraction

In Figure 1, the law of refraction applied at point  $P$  gives

$$(n + dn) \sin z = n \sin (z + dz) = n (\sin z + \cos z dz)$$

from which immediately follows the differential equation  $d(\Delta z) = dz = (\tan z/n) dn$  and the corresponding integral equation

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$$(0 \leq z_1 \leq 90^\circ) ; \quad \Delta z = \int_1^{n_1} \frac{\tan z}{z} \, dn \quad (1)$$

which is the basic mathematical expression of the correction for astronomical refraction.

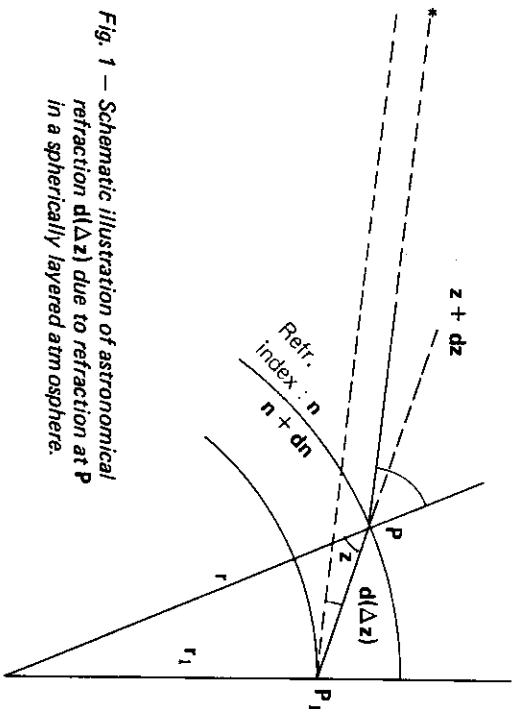


Fig. 1 — Schematic illustration of astronomical refraction  $d(\Delta z)$  due to refraction at  $P$  in a spherically layered atmosphere.

Since  $z$  is not, in general, constant along the light path but depends upon the refractive index according to the law of refraction

$$nr \sin z = n_1 r_1 \sin z_1 = \text{const.} \quad (2)$$

it will be necessary to find a suitable expression for  $\tan z$  that makes (1) integrable. Setting  $n_1 r_1 / (nr) = y$  for brevity, we have from (2)

$$\sin^2 z = y^2 \tan^2 z_1 / (1 + \tan^2 z_1)$$

$$\cos^2 z = (1 + \tan^2 z_1 - y^2 \tan^2 z_1) / (1 + \tan^2 z_1)$$

and

$$\tan z = y \tan z_1 [1 + \tan^2 z_1 (1 - y^2)]^{-\frac{1}{2}}$$

$$= y \tan z_1 - \frac{1}{2} y (1 - y^2) \tan^3 z_1 + \frac{3}{8} y (1 - y^2)^2 \tan^5 z_1 - \dots$$

$$- \frac{5}{16} y (1 - y^2)^3 \tan^7 z_1 + \frac{35}{128} y (1 - y^2)^4 \tan^9 z_1 - \frac{63}{256} y (1 - y^2)^5 \tan^{11} z_1 + \dots$$

Neglecting the subsequent terms in the binomial expansion, the first five may be written identically

$$\begin{aligned} \tan z = & \tan z_1 + \left( y - \frac{r_1}{r} \right) \left( \tan^3 z_1 + \tan z_1 \right) - \left( \frac{r - r_1}{r} \right)^2 \left( \tan^5 z_1 + \tan z_1 \right) + \\ & + \left( 1 + \frac{1}{2} y \right) (1 - y)^2 \tan^3 z_1 + \frac{3}{8} y (1 + y)^2 (1 - y)^2 \tan^5 z_1 - \\ & - \frac{5}{16} y (1 + y)^3 (1 - y)^3 \tan^7 z_1 + \frac{35}{128} y (1 + y)^4 (1 - y)^4 \tan^9 z_1 \end{aligned}$$

into which we substitute the approximation

$$y - \frac{r_1}{r} = \left( \frac{r_1}{r} \right) \left( \frac{n_1 - n}{n} \right) = n_1 - n$$

$$\frac{r - r_1}{r} = \frac{1}{r} (r - r_1) - \frac{1}{r_1} (r - r_1)^2$$

$$\left( 1 + \frac{1}{2} y \right) (1 - y)^2 = \frac{3}{2} (1 - y)^2 = \frac{3}{2 r_1^2} (r - r_1)^2$$

$$\frac{3}{8} y (1 + y)^3 (1 - y)^3 = \frac{3}{2} (1 - y)^2 = \frac{3}{2} \left[ \left( 1 - \frac{r_1}{r} \right) - (n_1 - n) \right]^2 =$$

$$= \frac{3}{2 r_1^2} (r - r_1)^2 - \frac{3}{r_1} (n_1 - n) (r - r_1)$$

$$\frac{5}{16} y (1 + y)^3 (1 - y)^3 = \frac{5}{2} (1 - y)^3 = \frac{5}{2 r_1^3} (r - r_1)^3$$

$$\frac{35}{128} y (1 + y)^4 (1 - y)^4 = \frac{35}{8} (1 - y)^4 = \frac{35}{8 r_1^4} (r - r_1)^4$$

and obtain

$$\begin{aligned} \tan z = & \tan z_1 + (\tan^3 z_1 + \tan z_1) (n_1 - n) - A_1 (r - r_1) + A_2 (r - r_1)^2 - \\ & - A_2 (n_1 - n) (r - r_1) - A_3 (r - r_1)^3 + A_4 (r - r_1)^4 \end{aligned} \quad (3)$$

where the coefficients are :

$$\begin{aligned}
 A_1 &= (\tan^3 z_1 + \tan z_1) / r_1 \\
 A_2 &= (3 \tan^5 z_1 + 5 \tan^3 z_1 + \dots) / (2 r_1^2) \\
 A_2' &= 3 \tan^5 z_1 / r_1 \\
 A_3 &= 5 \tan^7 z_1 / (2 r_1^3) \\
 A_4 &= 35 \tan^9 z_1 / (8 r_1^4)
 \end{aligned}
 \tag{4}$$

By the substitution of (3), integral (1) breaks down into seven terms, of which the first two can be solved at once :

$$\begin{aligned}
 \tan z_1 \int_1^{n_1} \frac{dn}{n} &= \tan z_1 \log n_1 = \tan z_1 \log [1 + (n_1 - 1)] = \\
 &= \tan z_1 (n_1 - 1) - \frac{1}{2} \tan z_1 (n_1 - 1)^2 + \dots \\
 (\tan^3 z_1 + \tan z_1) \int_1^{n_1} (n_1 - n) dn &= \frac{1}{2} (\tan^3 z_1 + \tan z_1) (n_1 - 1)^2
 \end{aligned}$$

Since  $n$  is nearly unity ( $1 \leq n < 1.0004$ ), all the terms of higher than second order will be omitted in the first integral, as well as  $n$  in the denominator of the subsequent ones. Equation (1) then becomes

$$\begin{aligned}
 \Delta z &= \tan z_1 (n_1 - 1) + \frac{1}{2} \tan^3 z_1 (n_1 - 1)^2 - A_1 \int_1^{n_1} (r - r_1) dn + \\
 &+ A_2 \int_1^{n_1} (r - r_1)^2 dn - A_2' \int_1^{n_1} (n_1 - n) (r - r_1) dn - \\
 &- A_3 \int_1^{n_1} (r - r_1)^3 dn + A_4 \int_1^{n_1} (r - r_1)^4 dn
 \end{aligned}
 \tag{5}$$

The five remaining atmospheric integrals can be determined, as follows.

$$\text{Integral} \int_1^{n_1} (r - r_1) dn.$$

In physical meteorology, the atmosphere may be thought of as a mixture of two ideal gases, dry air and water vapour. If we denote the total pressure, the partial pressure of water vapour and the absolute temperature by  $p$ ,  $e$  and  $T$  respectively, the densities of the dry-air and water-vapour components are, as stated by the perfect gas law.

$$\rho_d = \frac{p - e}{RT} \quad \text{and} \quad \rho_w = \frac{e}{R_w T}$$

where  $R$  and  $R_w$  stand for the appropriate gas constants. The density of the mixture is, of course, equal to  $\rho_d + \rho_w$ , or

$$\rho = \frac{p}{RT} - \left(1 - \frac{R}{R_w}\right) \frac{e}{RT}$$

The atmosphere being in hydrostatic equilibrium, pressure  $p$  measured at any height level is equal to the total weight of the air contained in a vertical column of unit cross section, reaching from the point of observation ( $r = r_1$ ) up to the top of the atmosphere ( $r = r'$ ). Consequently,

$$\int_{r_1}^{r'} \rho dr = \frac{1}{R} \int_{r_1}^{r'} \left(\frac{p}{T}\right) dr - \frac{1}{R} \left(1 - \frac{R}{R_w}\right) \int_{r_1}^{r'} \left(\frac{e}{T}\right) dr = \frac{p_1}{g} \tag{6}$$

where  $g$  is the local value of gravity at the centroid of the atmospheric column.

The refractivity of moist air for electromagnetic radiation may be written

$$n - 1 = \frac{(n_0 - 1) T_0}{p_0} \left(\frac{p}{T}\right) - c_w \left(\frac{e}{T}\right) + c_w' \left(\frac{e}{T^2}\right) \tag{7}$$

where  $n_0$  is the refractive index of dry air at pressure  $p_0$  and temperature  $T_0$ , and  $c_w$  and  $c_w'$  are constants. The corresponding height integral

$$\int_{r_1}^{r'} (n - 1) dr = \frac{(n_0 - 1) T_0}{p_0} \int_{r_1}^{r'} \left(\frac{p}{T}\right) dr - c_w \int_{r_1}^{r'} \left(\frac{e}{T}\right) dr + c_w' \int_{r_1}^{r'} \left(\frac{e}{T^2}\right) dr$$

can be readily determined with the aid of equation (6). This gives

$$\int_{r_1}^{r'} (n-1) dr = \frac{(n_0-1)RT_0}{p_0 g} p_1 + \left[ \frac{(n_0-1)T_0}{p_0} \left( 1 - \frac{R}{R_w} \right) - c_w \right] \int_{r_1}^{r'} \left( \frac{e}{T} \right) dr + c_w' \int_{r_1}^{r'} \left( \frac{e}{T^2} \right) dr \tag{8}$$

Equation (8) expresses the value of the refractivity integral in terms of ground pressure  $p_1$ , with minor corrections included due to the presence of water vapour in the atmosphere.

As far as the astronomical refraction is concerned, the contribution of humidity to the refractivity integral is negligible, and the last two terms in equation (8) can be omitted. Setting

$$u = r - r_1 \qquad v = n - 1$$

$$du = dr \qquad dv = dn$$

and integrating by parts :

$$\int (r - r_1) dn = uv - \int v du = (r - r_1) (n - 1) - \int (n - 1) dr,$$

we then obtain from (8) and (7) the important relationships

$$\int_{r_1}^{n_1} (r - r_1) dn = \int_{r_1}^{r'} (n - 1) dr = \frac{(n_0 - 1)RT_0}{p_0 g} p_1 = \frac{R}{g} (n_1 - 1)T_1 \tag{9}$$

$$\text{Integral} \int_{r_1}^{n_1} (r - r_1)^2 dn.$$

This integral requires some consideration of the vertical distribution of pressure and temperature in the atmosphere. We shall determine its value in two parts, the stratospheric component and the tropospheric component. The state of the atmosphere at the bounding surface, the tropopause, shall be denoted by superscripts  $p^0, T^0$ , etc., and it is assumed to be known.

Throughout the stratosphere, the temperature may be taken as constant, and equal to temperature  $T^0$  at the tropopause. Integration of the hydrostatic equation for fluids,  $dp = g \rho dr$ , on the condition  $\rho = p/(RT^0)$  gives the pressure as

$$p = p^0 e^{m(r - r^0)} \tag{10}$$

where  $e$  is the base of natural logarithms, and  $m = -g/(RT^0)$  is constant. Similarly,

$$n - 1 = (n^0 - 1) e^{m(r - r^0)} \tag{11}$$

and differentiating (11)

$$dn = m (n^0 - 1) e^{m(r - r^0)} dr = m (n - 1) dr \tag{12}$$

Since identically

$$r - r_1 = (r^0 - r_1) + (r - r^0)$$

$$(r - r_1)^2 = (r^0 - r_1)^2 + 2(r^0 - r_1)(r - r^0) + (r - r^0)^2$$

we have first, using (9)

$$\int_{r_1}^{n^0} (r - r_1)^2 dn = (r^0 - r_1)^2 (n^0 - 1) + \frac{2R}{g} (r^0 - r_1) (n^0 - 1)T^0 + \int_{r_1}^{n^0} (r - r^0)^2 dn$$

Now from (12)

$$\int (r - r^0)^2 dn = m (n^0 - 1) \int (r - r^0)^2 e^{m(r - r^0)} dr =$$

$$= m (n^0 - 1) \frac{e^{m(r - r^0)}}{m^3} \left[ m^2 (r - r^0)^2 - 2m (r - r^0) + 2 \right] + C =$$

$$= (n - 1) \left[ (r - r^0)^2 - \frac{2}{m} (r - r^0) + \frac{2}{m^2} \right] + C$$

where  $C$  is the constant of integration. This gives

$$\int_{r_1}^{n^0} (r - r^0)^2 dn = \frac{2(n^0 - 1)}{m^2} = \frac{2R^2}{g^2} (n^0 - 1)T^{02} \tag{13}$$

and the total stratospheric component is consequently

$$\int_{r_1}^{n^0} (r - r_1)^2 dn = (r^0 - r_1)^2 (n^0 - 1) + \frac{2R}{g} (r^0 - r_1) (n^0 - 1)T^0 + \frac{2R^2}{g^2} (n^0 - 1)T^{02} \tag{14}$$

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Through most of the troposphere, the temperature decreases with height at a fairly uniform rate which varies slightly with latitude and season, although in the polar regions there exists a permanent inversion in the lower troposphere where the actual temperatures increase with height. Integration of the hydrostatic equation on the conditions  $\rho = p/(RT)$  and

$$T = T_1 + \beta(r - r_1) \tag{15}$$

where the vertical gradient of temperature,  $\beta = dT/dr$ , is assumed constant gives the pressure as

$$p = p_1 \left( \frac{T}{T_1} \right)^{-g/(R\beta)} \tag{16}$$

and the pressure-temperature ratio as  $p/T = (p_1/T_1) (T/T_1)^{m'}$ , where  $m' = -g/(R\beta) - 1$  is constant. The refractivity is now given by

$$n - 1 = (n_1 - 1) \left( \frac{T}{T_1} \right)^{m'} \tag{17}$$

and its differential by

$$dn = \frac{m'(n_1 - 1)}{T_1} \left( \frac{T}{T_1} \right)^{m'-1} dT = m' \left( \frac{n - 1}{T} \right) dT \tag{18}$$

Since from (15)

$$r - r_1 = \frac{T - T_1}{\beta} = \frac{T_1}{\beta} \left( \frac{T}{T_1} - 1 \right) \quad \text{and} \quad (r - r_1)^2 = \frac{T_1^2}{\beta^2} \left( \frac{T}{T_1} - 1 \right)^2$$

we have

$$\begin{aligned} \int (r - r_1)^2 dn &= \frac{m'(n_1 - 1) T_1}{\beta^2} \int \left( \frac{T}{T_1} - 1 \right)^2 \left( \frac{T}{T_1} \right)^{m'-1} dT = \\ &= \frac{m'(n_1 - 1) T_1^2}{\beta^2} \left[ \left( \frac{1}{m'+2} \right) \left( \frac{T}{T_1} \right)^{m'+2} - \left( \frac{2}{m'+1} \right) \left( \frac{T}{T_1} \right)^{m'+1} + \frac{1}{m'} \left( \frac{T}{T_1} \right)^m \right] + C = \\ &= \frac{(n - 1) T_1^2}{\beta^2} \left[ \left( \frac{m'}{m'+2} \right) \left( \frac{T}{T_1} \right)^2 - \left( \frac{2m'}{m'+1} \right) \left( \frac{T}{T_1} \right) + 1 \right] + C = \end{aligned}$$

$$= (r - r_1)^2 (n - 1) + \frac{(n - 1) T_1^2}{\beta^2} \left[ \left( \frac{2}{m'+1} \right) \left( \frac{T}{T_1} \right) - \left( \frac{2}{m'+2} \right) \left( \frac{T}{T_1} \right)^2 \right] + C$$

where  $C$  is the constant of integration, and the preceding term, transformed step by step, is

$$\begin{aligned} \frac{(n - 1) T_1^2}{\beta^2} \left[ \left( \frac{2}{m'+1} \right) \left( \frac{T}{T_1} \right) - \left( \frac{2}{m'+2} \right) \left( \frac{T}{T_1} \right)^2 \right] &= \frac{2(n - 1) T_1 T}{\beta^2 (m'+1)} \left[ 1 - \left( \frac{m'+1}{m'+2} \right) \left( \frac{T}{T_1} \right) \right] = \\ &= \frac{2(n - 1) T_1 T}{\beta^2 (m'+1)} \left[ \left( \frac{1}{m'+2} \right) \left( \frac{T}{T_1} \right) - \left( \frac{T}{T_1} - 1 \right) \right] = \frac{2(n - 1) T}{\beta (m'+1)} \left[ \frac{T}{\beta (m'+2)} - (r - r_1) \right] = \\ &= \frac{2(n - 1) RT}{g} \left[ (r - r_1) - \frac{T}{\beta (m'+2)} \right] = \frac{2R}{g} (r - r_1) (n - 1) T + \frac{2R^2}{g^2 (1 - R\beta/g)} (n - 1) T^2 \end{aligned}$$

The tropospheric component is accordingly

$$\begin{aligned} \int_{n^0}^{n^1} (r - r_1)^2 dn &= -(r^0 - r_1)^2 (n^0 - 1) - \frac{2R}{g} (r^0 - r_1) (n^0 - 1) T^0 + \\ &+ \frac{2R^2}{g^2 (1 - R\beta/g)} \left[ (n_1 - 1) T_1^2 - (n^0 - 1) T^{02} \right] \tag{19} \end{aligned}$$

which equation holds for any constant value of  $\beta \neq g/R$ , including  $\beta = 0$ .

The sum of component integrals (14) and (19) gives the total value of the integral

$$\int_1^{n^1} (r - r_1)^2 dn = \frac{2R^2}{g^2} \left[ \frac{(n_1 - 1) T_1^2 - (n^0 - 1) T^{02}}{1 - R\beta/g} + (n^0 - 1) T^{02} \right] \tag{20}$$

under normal atmospheric conditions where the vertical distribution of temperature throughout the troposphere is substantially a linear function of height.

$$\text{Integral} \int_1^{n^1} (n_1 - n) (r - r_1) dn.$$

Integration by parts using the substitutions  $u = r - r_1$  and  $v = [n_1 - 1 - (n - 1)]^2$  gives first, in view of (9),

$$\int_1^{n_1} (n_1 - n) (r - r_1) \, dn = \frac{R}{g} (n_1 - 1)^2 T_1 - \frac{1}{2} \int_{r_1}^{r'} (n - 1)^2 \, dr \quad (21)$$

the latter integral being more conveniently determined.

In the stratosphere, equation (11) gives

$$\begin{aligned} \int (n - 1)^2 \, dr &= (n^0 - 1)^2 \int e^{2m(r - r^0)} \, dr = \\ &= (n^0 - 1)^2 \frac{e^{2m(r - r^0)}}{2m} + C = -\frac{R}{2g} (n - 1)^2 T^0 + C \end{aligned}$$

and

$$\int_0^{r'} (n - 1)^2 \, dr = \frac{R}{2g} (n^0 - 1)^2 T^0 \quad (22)$$

whereas in the troposphere, applying (17)

$$\begin{aligned} \int (n - 1)^2 \, dr &= \frac{(n_1 - 1)^2}{\beta} \int \left( \frac{T}{T_1} \right)^{2m'} \, dT = \\ &= \frac{(n_1 - 1)^2 T_1}{\beta (2m' + 1)} \left( \frac{T}{T_1} \right)^{2m' + 1} + C = -\left( \frac{R}{2g + R\beta} \right) (n - 1)^2 T + C \end{aligned}$$

and

$$\int_{r_1}^{r^0} (n - 1)^2 \, dr = \left( \frac{R}{2g + R\beta} \right) [ (n_1 - 1)^2 T_1 - (n^0 - 1)^2 T^0 ] \quad (23)$$

The sum of (22) and (23) substituted into equation (21) finally gives the total value of the integral

$$\begin{aligned} \int_1^{n_1} (n_1 - n) (r - r_1) \, dn &= \frac{R}{g} (n_1 - 1)^2 T_1 - \frac{R}{2(2g + R\beta)} [ (n_1 - 1)^2 T_1 + \\ &+ \frac{1}{2} (R\beta/g) (n^0 - 1)^2 T^0 ] \quad (24) \end{aligned}$$

again assuming that the vertical gradient of temperature is constant in the troposphere.

$$\text{Integrals } \int_1^{n_1} (r - r_1)^3 \, dn \text{ and } \int_1^{n_1} (r - r_1)^4 \, dn.$$

For the stratospheric component of the first integral we have from (12)

$$\begin{aligned} \int (r - r^0)^3 \, dn &= m (n^0 - 1) \int (r - r^0)^3 e^{m(r - r^0)} \, dr = \\ &= m (n^0 - 1) \left[ \frac{1}{m} (r - r^0)^3 e^{m(r - r^0)} - \frac{3}{m} \int (r - r^0)^2 e^{m(r - r^0)} \, dr \right] = \\ &= (r - r^0)^3 (n - 1) - \frac{3}{m} \int (r - r^0)^2 \, dn \end{aligned}$$

and further, in view of (13)

$$\int_1^{n^0} (r - r^0)^3 \, dn = -\frac{3}{m} \int_1^{n^0} (r - r^0)^2 \, dn = \frac{6R^3}{g^3} (n^0 - 1) T^{03} \quad (25)$$

Using the identity  $(r - r_1)^3 = (r^0 - r_1)^3 + 3(r^0 - r_1)^2 (r - r^0) + 3(r^0 - r_1) (r - r^0)^2 + (r - r^0)^3$  and applying integrals (9), (13), and (25), the stratospheric component is obtained as

$$\begin{aligned} \int_1^{n^0} (r - r_1)^3 \, dn &= (r^0 - r_1)^3 (n^0 - 1) + \frac{3R}{g} (r^0 - r_1)^2 (n^0 - 1) T^0 + \\ &+ \frac{6R^2}{g^2} (r^0 - r_1) (n^0 - 1) T^{02} + \frac{6R^3}{g^3} (n^0 - 1) T^{03} \quad (26) \end{aligned}$$

For the tropospheric component of the same integral we have from (15) and (18)

$$\begin{aligned} \int (r - r_1)^3 \, dn &= \frac{m'(n_1 - 1) T_1^2}{\beta^3} \int \left( \frac{T}{T_1} - 1 \right)^3 \left( \frac{T}{T_1} \right)^{m' - 1} \, dT = \\ &= \frac{m'(n_1 - 1) T_1^3}{\beta^3} \left[ \left( \frac{1}{m' + 3} \right) \left( \frac{T}{T_1} \right)^{m' + 3} - \left( \frac{3}{m' + 2} \right) \left( \frac{T}{T_1} \right)^{m' + 2} + \left( \frac{3}{m' + 1} \right) \left( \frac{T}{T_1} \right)^{m' + 1} - \right. \\ &\left. - \frac{1}{m'} \left( \frac{T}{T_1} \right)^{m'} \right] + C = \frac{(n - 1) T_1^3}{\beta^3} \left[ \left( \frac{m'}{m' + 3} \right) \left( \frac{T}{T_1} \right)^3 - \left( \frac{3m'}{m' + 2} \right) \left( \frac{T}{T_1} \right)^2 + \left( \frac{3m'}{m' + 1} \right) \left( \frac{T}{T_1} \right) - 1 \right] + C = \end{aligned}$$

$$\begin{aligned}
 &= (r-r_1)^3 (n-1) - \frac{3(n-1)T_1^2 T}{\beta^3 (m'+1)} \left[ \left( \frac{m'+1}{m'+3} \right) \left( \frac{T}{T_1} \right)^2 - \left( \frac{2m'+2}{m'+2} \right) \left( \frac{T}{T_1} \right) + 1 \right] + C = \\
 &= (r-r_1)^3 (n-1) + \frac{3R}{g} (r-r_1)^2 (n-1) T + \frac{6(n-1)T_1 T^2}{\beta^3 (m'+1)(m'+2)} \left[ \left( \frac{m'+2}{m'+3} \right) \left( \frac{T}{T_1} \right) - 1 \right] + C \\
 &= (r-r_1)^3 (n-1) + \frac{3R}{g} (r-r_1)^2 (n-1) T + \frac{6R^2}{g^2 (1-R\beta/g)} (r-r_1) (n-1) T^2 + \\
 &+ \frac{6R^3}{g^3 (1-R\beta/g)(1-2R\beta/g)} (n-1) T^3 + C
 \end{aligned}$$

The tropospheric component is accordingly

$$\begin{aligned}
 \int_0^{n_1} (r-r_1)^3 dn &= -(r^0-r_1)^3 (n^0-1) - \frac{3R}{g} (r^0-r_1)^2 (n^0-1) T^0 - \\
 &- \frac{6R^2}{g^2 (1-R\beta/g)} (r^0-r_1) (n^0-1) T^{02} + \\
 &+ \frac{6R^3}{g^3 (1-R\beta/g)(1-2R\beta/g)} [(n_1-1)T_1^3 - (n^0-1)T^{03}]
 \end{aligned} \tag{27}$$

which added to stratospheric component (26) gives the total integral

$$\int_1^{n_1} (r-r_1)^3 dn = \frac{6R^3}{g^3} \left[ \frac{(n_1-1)T_1^3 - (n^0-1)T^{03}}{(1-R\beta/g)(1-2R\beta/g)} + (n^0-1)T^{03} \right] + \tag{28}$$

$$+ \frac{6R^2}{g^2} \left[ 1 - \frac{1}{1-R\beta/g} \right] (r^0-r_1) (n^0-1) T^{02}$$

Similarly,

$$\begin{aligned}
 \int_1^{n_1} (r-r_1)^4 dn &= \frac{24R^4}{g^4} \left[ \frac{(n_1-1)T_1^4 - (n^0-1)T^{04}}{(1-R\beta/g)(1-2R\beta/g)(1-3R\beta/g)} + (n^0-1)T^{04} \right] + \\
 &+ \frac{24R^3}{g^3} \left[ 1 - \frac{1}{(1-R\beta/g)(1-2R\beta/g)} \right] (r^0-r_1) (n^0-1) T^{03} + \\
 &+ \frac{12R^2}{g^2} \left( 1 - \frac{1}{1-R\beta/g} \right) (r^0-r_1)^2 (n^0-1) T^{02}
 \end{aligned} \tag{29}$$

is obtained for the explicit value of the last integral considered in the expression for astronomical refraction (5).

Due to insolational heating of the ground during the day and its radiational cooling during the night, temperature gradients within the first few kilometres of the troposphere next to the ground frequently differ significantly from the approximately constant value of  $\beta$  above that level. Consequently integrals (20), (24), (28) and (29) should be modified by subdividing their respective tropospheric components. But since only a small contribution to these integrals comes from the lower levels, it will be quite sufficient merely to extend the constant temperature gradient of the free troposphere down to the ground level, neglecting the small error thus involved. This requires that the actual values of  $T_1$  and  $n_1 - 1$  be replaced by

$$T_1' = T^0 - \beta(r^0 - r_1) \tag{15'}$$

$$n_1' - 1 = (n^0 - 1) (T_1' / T^0)^m \tag{17'}$$

and the prevailing temperature gradient of the lower troposphere can be disregarded.

We may now combine the results from the preceding discussion, and write on the basis of equation (5) the following expression for the correction for astronomical refraction (in seconds of arc) :

$$\begin{aligned}
 \Delta z'' &= \rho'' \tan z_1 \left[ 1 + \frac{1}{2} \tan^2 z_1 (n_1 - 1) \right] (n_1 - 1) - \\
 &- \frac{\rho'' R}{r_1 g} (\tan^3 z_1 + \tan z_1) (n_1 - 1) T_1 + \delta_1'' - \delta_2'' - \delta_3'' + \delta_4''
 \end{aligned} \tag{30}$$

where

$$\begin{aligned}
 \delta_1'' &= \frac{\rho'' R^2}{r_1^2 g^2} (3 \tan^5 z_1 + 5 \tan^3 z_1) \left[ \frac{(n_1' - 1) T_1'^2 - (n^0 - 1) T^{02}}{1 - R\beta/g} + (n^0 - 1) T^{02} \right] \\
 \delta_2'' &= \frac{3 \rho'' R}{r_1 g} \tan^5 z_1 \left[ (n_1' - 1)^2 T_1' - \frac{(n_1' - 1)^2 T_1' + \frac{1}{2} (R\beta/g) (n^0 - 1)^2 T^0}{2(2 + R\beta/g)} \right] \\
 \delta_3'' &= \frac{15 \rho'' R^3}{r_1^3 g^3} \tan^7 z_1 \left[ \frac{(n_1' - 1) T_1'^3 - (n^0 - 1) T^{03}}{(1 - R\beta/g)(1 - 2R\beta/g)} + (n^0 - 1) T^{03} \right] + \\
 &+ \frac{15 \rho'' R^2}{r_1^3 g^2} \tan^7 z_1 \left( 1 - \frac{1}{1 - R\beta/g} \right) (r^0 - r_1) (n^0 - 1) T^{02}
 \end{aligned} \tag{31}$$

$$\delta_4'' = \frac{105 \rho'' R^4}{r_1^4 g^4} \tan^9 z_1 \left[ \frac{(n_1' - 1) T_1'^4 - (n^0 - 1) T^{04}}{(1 - R\beta/g)(1 - 2R\beta/g)(1 - 3R\beta/g)} + (n^0 - 1) T^{04} \right] +$$

$$+ \frac{105 \rho'' R^3}{r_1^4 g^3} \tan^9 z_1 \left[ \frac{1}{(1 - R\beta/g)(1 - 2R\beta/g)} \right] (r^0 - r_1) (n^0 - 1) T^{03} +$$

$$+ \frac{105 \rho'' R^2}{2 r_1^4 g^2} \tan^9 z_1 \left( 1 - \frac{1}{1 - R\beta/g} \right) (r^0 - r_1)^2 (n^0 - 1) T^{02}$$

represent minor terms dependent on the vertical structure of the atmosphere. Up to zenith distance  $z_1 = 80^\circ$ , equation (30) will give the value of integral (1) accurately enough for all practical purposes, as can best be demonstrated by test computations on atmospheric models based upon the formulas previously derived (see Tables 1a - c and 11a - c).

Table 1a.  
Atmospheric Model No. 1  
(Tropical Zone)

r, km 6360+	p, mb	T, °K	(n-1)10 <sup>6</sup>	-(dn/dr) 10 <sup>6</sup> , km <sup>-1</sup>	r, km 6360+	p, mb	T, °K	(n-1) 10 <sup>6</sup>	-(dn/dr) 10 <sup>6</sup> , km <sup>-1</sup>
0	1010.00	299.85	265.72	24.8210	21.3	45.19	198.00	18.01	3.0983
0.8	921.55	295.00	246.43	23.3982	22.2	38.71	198.00	15.42	2.6538
1.6	839.57	290.15	228.26	22.0353	23.1	33.15	198.00	13.21	2.2730
2.4	763.68	285.30	211.16	20.7308	24	28.40	198.00	11.31	1.9469
3.2	693.53	280.45	195.08	19.4832	26	20.13	198.00	8.02	1.3800
4	628.76	275.60	179.97	18.2908	28	14.27	198.00	5.68	0.9782
4.8	569.05	270.75	165.80	17.1522	30	10.11	198.00	4.03	0.6933
5.6	514.08	265.90	152.52	16.0658	32	7.17	198.00	2.86	0.4914
6.4	463.55	261.05	140.08	15.0300	34	5.08	198.00	2.02	0.3483
7.7	390.19	253.17	121.58	13.4513	36	3.60	198.00	1.43	0.2469
9	326.65	245.29	105.05	11.9962	38	2.55	198.00	1.02	0.1750
10.3	271.88	237.41	90.34	10.6585	40	1.81	198.00	0.72	0.1241
11.6	224.89	229.52	77.29	9.4323	44	0.91	198.00	0.36	0.0623
12.9	184.80	221.64	65.77	8.3116	48	0.46	198.00	0.18	0.0313
14.2	150.77	213.76	55.64	7.2906	52	0.23	198.00	0.09	0.0157
15.5	122.08	205.88	46.78	6.3636	56	0.12	198.00	0.05	0.0079
16.8	98.03	198.00	39.06	5.5250	60	0.06	198.00	0.02	0.0040
16.8	98.03	198.00	39.06	6.7209	64	0.03	198.00	0.01	0.0020
17.7	83.97	198.00	33.45	5.7566	68	0.01	198.00	0.006	0.0010
18.6	71.92	198.00	28.65	4.9307	72	0.007	198.00	0.003	0.0005
19.5	61.60	198.00	24.54	4.2233					
20.4	52.76	198.00	21.02	3.6173					

$r_1 = 6360$  km  
 $r^0 = 6376.8$  km  
 $p_1 = 1010$  mb  
 $T_1 = 299.85$  °K  
 $n_1 = 1.000265717$  ( $\lambda = 0.574 \mu$ )  
 $\beta = -6.0625$  °K km<sup>-1</sup>  
 $R = 2.8704 \times 10^6$  erg g<sup>-1</sup> °K<sup>-1</sup>  
 $g = 97.8 \times 10^1$  cm sec<sup>-2</sup>



$r, \text{ km}$	$T, \text{ }^\circ\text{K}$	$p, \text{ mb}$	$1020.00$	$318.67$	$56.9646$	$18.3$	$70.68$	$223.00$	$25.00$	$3.8358$
0	1020.00	252.50	318.67	56.9646	18.3	70.68	223.00	25.00	3.8358	
0.2	992.85	254.68	307.53	54.5009	20.2	52.81	223.00	18.68	2.8659	
0.4	966.64	256.87	296.66	52.1633	22.1	39.46	223.00	13.96	2.1413	
0.6	941.34	261.24	276.88	49.9446	24	29.48	223.00	10.43	1.5999	
0.8	916.90	261.24	276.88	47.8378	26	21.69	223.00	7.67	1.1771	
1	893.30	263.42	267.51	45.8362	28	15.96	223.00	5.65	0.8661	
1.2	870.48	265.61	258.53	43.9339	30	11.74	223.00	4.15	0.6373	
1.4	848.44	267.80	249.93	42.1252	32	8.64	223.00	3.06	0.4689	
1.6	827.12	269.98	241.68	40.4047	34	6.36	223.00	2.25	0.3450	
1.8	807.12	272.12	233.98	38.7840	36	4.68	223.00	1.65	0.2538	
2.0	787.04	274.11	226.15	37.2779	38	3.44	223.00	1.22	0.1868	
2.5	737.04	278.24	200.11	33.0779	40	2.53	223.00	0.90	0.1374	
3.0	687.04	282.24	181.50	29.9125	44	1.37	223.00	0.48	0.0744	
4.0	587.04	292.36	148.27	18.4486	48	0.74	223.00	0.26	0.0403	
5.0	487.04	298.52	119.91	14.5050	52	0.40	223.00	0.14	0.0218	
6.0	387.04	299.39	107.39	13.3324	56	0.22	223.00	0.08	0.0118	
7.0	307.04	299.87	107.39	13.3324	60	0.12	223.00	0.04	0.0064	
8.0	247.04	299.87	107.39	13.3324	64	0.06	223.00	0.02	0.0035	
8.8	207.04	299.87	107.39	13.3324	68	0.03	223.00	0.01	0.0019	
10.0	167.04	299.87	107.39	13.3324	72	0.02	223.00	0.007	0.0010	
10.7	147.04	299.87	107.39	13.3324						
12.6	107.04	299.87	107.39	13.3324						
14.5	67.04	299.87	107.39	13.3324						
16.4	27.04	299.87	107.39	13.3324						

$r_1 = 6400 \text{ km}$   
 $r_0 = 6401.6 \text{ km}$   
 $r_0 = 6408.8 \text{ km}$   
 $p_1 = 1020 \text{ mb}$   
 $T_1 = 252.5 \text{ }^\circ\text{K}$   
 $n_1 = 1.000318670 (\lambda = 0.574 \mu)$   
 $\beta^1 = +10.925 \text{ }^\circ\text{K km}^{-1}$   
 $\beta = -6.525 \text{ }^\circ\text{K km}^{-1}$   
 $R = 2.8704 \times 10^6 \text{ erg g}^{-1} \text{ }^\circ\text{K}^{-1}$   
 $g = 98.2 \times 10^1 \text{ cm sec}^{-2}$

Table 1c. Atmospheric Model No. 3 (Arctic Zone)

$r, \text{ km}$	$T, \text{ }^\circ\text{K}$	$p, \text{ mb}$	$27.2824$	$26.2791$	$25.3018$	$24.3501$	$23.4236$	$22.5219$	$21.6447$	$20.7916$	$19.9622$	$18.6834$	$17.4629$	$16.2990$	$15.1902$	$14.1351$	$13.1322$	$12.1799$	$11.2767$	$13.9033$	$10.6535$	$8.1633$	$6.2551$	$4.7930$
0	285.08	280.87	27.2824	26.2791	25.3018	24.3501	23.4236	22.5219	21.6447	20.7916	19.9622	18.6834	17.4629	16.2990	15.1902	14.1351	13.1322	12.1799	11.2767	13.9033	10.6535	8.1633	6.2551	4.7930
0.5	955.68	281.86	267.48	25.3018	24.3501	23.4236	22.5219	21.6447	20.7916	19.9622	18.6834	17.4629	16.2990	15.1902	14.1351	13.1322	12.1799	11.2767	10.6535	8.1633	6.2551	4.7930		
1	899.20	278.63	254.58	24.3501	23.4236	22.5219	21.6447	20.7916	19.9622	18.6834	17.4629	16.2990	15.1902	14.1351	13.1322	12.1799	11.2767	10.6535	8.1633	6.2551	4.7930			
1.5	845.46	275.40	242.17	23.4236	22.5219	21.6447	20.7916	19.9622	18.6834	17.4629	16.2990	15.1902	14.1351	13.1322	12.1799	11.2767	10.6535	8.1633	6.2551	4.7930				
2	794.36	272.18	230.23	22.5219	21.6447	20.7916	19.9622	18.6834	17.4629	16.2990	15.1902	14.1351	13.1322	12.1799	11.2767	10.6535	8.1633	6.2551	4.7930					
2.5	745.79	268.96	218.74	21.6447	20.7916	19.9622	18.6834	17.4629	16.2990	15.1902	14.1351	13.1322	12.1799	11.2767	10.6535	8.1633	6.2551	4.7930						
3	699.66	265.73	207.70	20.7916	19.9622	18.6834	17.4629	16.2990	15.1902	14.1351	13.1322	12.1799	11.2767	10.6535	8.1633	6.2551	4.7930							
3.5	655.86	262.50	197.10	19.9622	18.6834	17.4629	16.2990	15.1902	14.1351	13.1322	12.1799	11.2767	10.6535	8.1633	6.2551	4.7930								
4	614.32	259.28	186.91	18.6834	17.4629	16.2990	15.1902	14.1351	13.1322	12.1799	11.2767	10.6535	8.1633	6.2551	4.7930									
4.8	552.31	254.12	171.45	17.4629	16.2990	15.1902	14.1351	13.1322	12.1799	11.2767	10.6535	8.1633	6.2551	4.7930										
5.6	495.48	248.96	157.00	16.2990	15.1902	14.1351	13.1322	12.1799	11.2767	10.6535	8.1633	6.2551	4.7930											
6.4	443.48	243.80	143.50	15.1902	14.1351	13.1322	12.1799	11.2767	10.6535	8.1633	6.2551	4.7930												
7.2	396.01	238.64	130.91	14.1351	13.1322	12.1799	11.2767	10.6535	8.1633	6.2551	4.7930													
8	352.74	233.48	119.18	13.1322	12.1799	11.2767	10.6535	8.1633	6.2551	4.7930														
8.8	313.38	228.32	108.28	12.1799	11.2767	10.6535	8.1633	6.2551	4.7930															
9.6	277.67	223.16	98.15	11.2767	10.6535	8.1633	6.2551	4.7930																
10.4	245.33	218.00	88.78	10.6535	8.1633	6.2551	4.7930																	
10.4	245.33	218.00	88.78	10.6535	8.1633	6.2551	4.7930																	
12.1	187.98	218.00	68.02	8.1633	6.2551	4.7930																		
13.8	144.04	218.00	52.12	6.2551	4.7930																			
15.5	110.37	218.00	39.94	4.7930																				
17.2	84.57	218.00	30.60																					

$r_1 = 6380 \text{ km}$   
 $r_0 = 6390.4 \text{ km}$   
 $p_1 = 1015 \text{ mb}$   
 $T_1 = 285.08 \text{ }^\circ\text{K}$   
 $n_1 = 1.000280868 (\lambda = 0.574 \mu)$   
 $\beta = -6.45 \text{ }^\circ\text{K km}^{-1}$   
 $R = 2.8704 \times 10^6 \text{ erg g}^{-1} \text{ }^\circ\text{K}^{-1}$   
 $g = 98 \times 10^1 \text{ cm sec}^{-2}$

Table 1b. Atmospheric Model No. 2 (Temperate Zone)

Table IIa. Atmospheric Model No. 1 (Tropical Zone)

Astronomical Refraction, Δz'' = -ρ'' ∫ (dn/dr) tan z dr, for z = 60°, 70° & 80°

Table with columns for r, km (6360+), -ρ'' (dn/dr/n), tan z (z1=60, 70, 80), -ρ'' (dn/dr/n) tan z (z1=60, 70, 80), r, km (6360+), -ρ'' (dn/dr/n), tan z (z1=60, 70, 80), -ρ'' (dn/dr/n) tan z (z1=60, 70, 80). Rows range from 0 to 20.4 km.

Integrals \*: 0 - 6.4 km: 44.8045, 70.9438, 145.0468; 6.4 - 16.8 km: 35.8720, 56.5197, 112.6666; 16.8 - 24 km: 9.8000, 15.3515, 29.7592; 24 - 40 km: 3.7214, 5.7954, 10.9443; 40 - 72 km: 0.2505, 0.3853, 0.6957

Formula (30): 1st term: 94.968, 150.735, 312.160; 2nd term: -0.525, -1.781, -14.264; δ1: 0.007, 0.052, 1.668; -δ2: -0.001, -0.007, -0.257; -δ3: -0.000, -0.002, -0.271; δ4: 0.000, 0.000, 0.061

\* Integration formula (Newton-Cotes):

∫\_a^a+BΔ f(r) dr = BΔ / 2.835 { 0.0989 [ f(a) + f(a+BΔ) ] + 0.5888 [ f(a+Δ) + f(a+7Δ) ] - 0.0928 [ f(a+2Δ) + f(a+6Δ) ] + 1.0496 [ f(a+3Δ) + f(a+5Δ) ] - 0.454 f(a+4Δ) }

Table IIb. Atmospheric Model No. 2 (Temperate Zone)

Astronomical Refraction, Δz'' = -ρ'' ∫ (dn/dr) tan z dr, for z = 60°, 70° & 80°

Table with columns for r, km (6380+), -ρ'' (dn/dr/n), tan z (z1=60, 70, 80), -ρ'' (dn/dr/n) tan z (z1=60, 70, 80), r, km (6380+), -ρ'' (dn/dr/n), tan z (z1=60, 70, 80), -ρ'' (dn/dr/n) tan z (z1=60, 70, 80). Rows range from 0 to 17.2 km.

Integrals \*: 0 - 4 km: 33.5267, 53.1215, 108.9898; 4 - 10.4 km: 34.9234, 56.1647, 111.3757; 10.4 - 24 km: 27.7120, 43.5495, 85.7216; 24 - 40 km: 3.4041, 5.3011, 10.0983; 40 - 72 km: 0.2969, 0.4591, 0.8274

Formula (30): 1st term: 100.386, 159.339, 330.039; 2nd term: -0.525, -1.781, -14.260; δ1: 0.006, 0.050, 1.599; -δ2: -0.001, -0.007, -0.272; -δ3: -0.000, -0.002, -0.258; δ4: 0.000, 0.000, 0.059

\* - See footnote to Table IIa.

Astronomical Refraction,  $\Delta z'' = -p'' \int_{z_1}^{z_2} \frac{dn}{dr} \tan z \, dr$ , for  $z = 60^\circ, 70^\circ, 80^\circ$

Table IIc.  
Atmospheric Model No. 3  
(Arctic Zone)

Main data table with columns for height (r, km), refraction (p'', arcmin), and astronomical refraction (Delta z'', arcmin) for various zenith angles (z1, z2) and distances.

Integrals:  
0 - 1.6 km: 27.4896 43.5899 89.8025  
1.6 - 8.8 km: 47.8525 75.7050 153.9958  
8.8 - 24 km: 34.3796 54.0825 106.9880  
24 - 40 km: 3.3500 5.2165 9.8585  
40 - 72 km: 0.3092 0.4769 0.8610

Formulae (30):  
1st term: 113.903 180.810 374.686  
2nd term: -0.525 -1.780 -14.257  
3rd term: 0.006 0.048 1.561  
4th term: -0.001 -0.009 -0.273  
5th term: -0.000 -0.002 -0.252  
6th term: 0.000 0.000 0.058

See footnote to Table IIa.

Summary table with columns for distance (km) and corresponding refraction values.

